

Spatio-Temporal Analysis of Climate Change Impact on Future Wind Power Potential in Burundi (East Africa)

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Abstract

This paper assessed climate change impact on future wind power potential across highlands and western lowlands of Burundi. Hourly observed and MERRA-2 data were considered for the historical period 1980-2016, and a Multi-model ensemble for future projections data of eight selected Regional Climate Models under RCP 4.5 and RCP 8.5 over the periods 2019-2040 and 2071-2100 was used. Variability and trend analysis were adopted using standardized index and Mann-Kendall's test, respectively while wind power density (WPD) in quartiles was adopted for changes distribution. As results, diurnal wind speeds (WS) were higher from 9:00 AM to 2:00 PM, while monthly wind speeds reached the maximum during summer time. An increasing trend in WPD was detected all over the studied area. Over the period 2019-2040, the lowest WPD change is projected at Northern Highlands (NHL) under RCP 4.5 with 28.04 W·m⁻² while the highest WPD change of 47.35 $W \cdot m^{-2}$ is projected under RCP 8.5 at Southern Imbo plain (SIP). As for the period 2071-2100, the highest change is expected at SIP under RCP 8.5 with 152.39 $W \cdot m^{-2}$ while the minimum change of 83.96 $W \cdot m^{-2}$ is projected under RCP 4.5 at NHL. The findings showed that areas nearby the Lake Tanganyika are expected to have high positive WDP changes.

Keywords

Burundi, Wind Power, Changes, Future Climate

1. Introduction

Climate influences the world through changing temperature, precipitation, snowmelt, and a host of other natural phenomena [1]. Effects of climate change

vary from one region to another depending on various socio-economic and environmental factors [2]. Globally, mean surface air temperature has risen by about $0.74^{\circ}C \pm 0.18^{\circ}C$ during the twentieth century and is projected to rise by $1.8^{\circ}C - 4.0^{\circ}C$ in the twenty-first century [3]. Extreme precipitations are also projected to increase significantly, especially in regions that are already wet under the current climate conditions, whereas dry spells are predicted to increase particularly in regions characterized by dry weather conditions in the current climate [2]. Furthermore, climate change is predicted to cause stronger surface wind speed values across the boreal regions of the Northern Hemisphere, including much of Canada, Siberia and northern Europe, and in tropical and subtropical regions in Africa, and Central and South America. However, Greenland, southern Europe, China, India, southern Australia and much of the west coast of South America are expected to experience decreasing wind speed values [4].

In energy sector, to mitigate climate change impacts, renewable sources have received considerable attention in power generation due to the growing public concern about the environmental degradation caused by conventional electricity generation and also due to the urgent need in exploring the possibilities of using alternative energy sources [5]. Among the various alternative sources of energy, the wind energy is the fastest growing and most promising renewable energy source as it is proved to be economically viable and commercialized [6] [7]. It does not pollute the environment and it is an endless and a free primary source. Therefore, the increasing demand for renewable energy sources in the coming decades requires that we have a clear understanding of accessible wind resources, and the susceptibility of these resources to climate change [8] [9].

In East Africa, wind energy, which remains hitherto untapped, appears promising in solving the problem of electricity shortage. For example, in Tanzania, Kainkwa [10] reported that the wind energy available in some locations in the northern part of the country was sufficient to generate electricity. Therefore, the acute shortage of electricity that prevails during the dry season could be alleviated to a large extent. Moreover, in Kenya, Peter [11] reported that Kenya has a proven wind energy potential of as high as 346 W·m⁻² and wind speeds of 6 m·s⁻¹ and above. Thus, the implementation of wind energy projects would significantly reduce the country's dependence on hydropower generation (around 50%) which drops quite significantly during dry seasons.

Burundi is an East African country heavily dependent on hydropower for his electricity production and experiences power shortage during the dry seasons [12]. To supplement the shortfall of electricity fossil fuels have become an alternative solution [13]. In fact, the fossil fuels produce a lot of harmful emissions to our environment. Thus, renewable energy sources such as wind, if thoroughly investigated, could be used to a large extent to reduce the dependence of fossil fuels in electricity generation during the dry seasons. Unfortunately, in Burundi, there are no many studies conducted to analyze projected wind power potential. However, according to Collins *et al.* [14], the climate change models predict that Burundi would experience more extreme weather conditions in the future cha-

racterized by high winds, hail, and extended droughts. Furthermore, Lawin *et al.* [15] detected increase in projected wind speed over the period 2021-2050 at Bujumbura station. But, neither of the studies analyzed projected wind power potential over the highlands and western lowlands. Therefore, this paper aims to revisit projected wind speeds and analyze impacts of wind speeds changes on projected wind power potential by considering four sites from two contrasting regions; western lowlands and highlands of Burundi. It fills then the gap by investigating changes in projections over near future (2019-2040) and far future (2071-2100) across highlands and western lowlands of Burundi where wind energy has not yet been extensively analyzed to date. Specifically, the study focuses on analyzing projected changes in wind speeds and evaluates their implication on projected wind power potential across Northern Highlands (NHL), Northern Imbo Plain (NIP), Southern Highlands (SHL) and Southern Imbo Plain (SIP).

2. Materials and Methods

2.1. Study Area

Figure 1 shows the study area location in Burundi, an East African country





between longitudes 28.8°E and 30.9°E and latitudes 2.3°S and 4.45°S [16]. Bounded to the north by Rwanda, to the East and South by Tanzania and to the West by the Democratic Republic of the Congo (DRC), Burundi covers an area of 27,834 km² [17]. The study area elevation is between 1800 and 2650 m altitude in the highlands while lowlands are between 774 and 1100 m altitude. The annual rainfall in the high altitude regions is almost the double that of low altitudes. Average maximum annual temperatures range from 21.8°C to 29.5°C. Burundi has average wind speeds varying between 4 m·s⁻¹ and 6 m·s⁻¹ [18]. The study regions are highlands and western lowlands of Burundi. For better analysis, the highlands were divided into Northern Highlands (NHL) and Southern Highlands (SHL) while the western lowlands were split into Northern Imbo Plain (NIP) and Southern Imbo Plain (SIP).

2.2. Data Used

Three sets of data are used in this study. The first set consists of observed data collected from the synoptic stations of Burundi belonging to the Geographical Institute of Burundi (IGEBU). In fact, nineteen meteorological stations have been considered over the historical period 1980-2016 and the concise information on their location are presented in **Figure 1**. Wind speed observed data were collected at hourly scales. The selection of the time series data was based on the quality of the data collected from various stations. However, only the station of Bujumbura (Airport) was able to cover the considered period 1980-2016, a period which respects the normal of 30 years at least, recommended by World Meteorology Organization (WMO).

To complement observed values at hourly scales, the second set of data grouping MERRA-2 (Modern Era Reanalysis for Research and Applications Version 2) data available online at

<u>https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/</u> was used. This data set is very useful in view of the fact that weather stations are limited in number, unevenly distributed, have missing data problem and short period of observation [19]. Missing observed data have been filled using cross validation method following Ioannis [20]. Table 1 presents the statistical summary of the used data at annual scale.

The third set of used data groups wind speeds from eight Regional Climate Models. These data are available in the context of the COordinated Regional

Tabl	e 1. Desci	ription o	f used d	ata over	the period	l 1980-2016 a	t annual sc	ale
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Variable	Region	Minimum	Maximum	Mean	Standard deviation
WS ⁺	NHL	3.8	4.8	4.2	0.216
	NIP	3.6	4.7	4.1	0.221
	SHL	3.8	4.9	4.3	0.231
	SIP	4.3	5.3	4.8	0.221

⁺WS = Wind speed in $(m \cdot s^{-1})$.

Climate Downscaling Experiment (CORDEX) over Africa at 0.44° resolution for the period 1950 to 2100 [21] and accessed online (https://www.cordex.org) through user registration. Table 2 shows the climatic models used to generate the multi-model ensemble, where the third column gives their acronyms adopted in this paper. Two experiments performed following the stabilization and the most extreme IPCC Representative Concentration Pathways scenarios (RCP 4.5 and RCP 8.5) for the periods 2019-2040 and 2071-2100 in CORDEX database have been considered. RCP 4.5 is a scenario that stabilizes radiative forcing at 4.5 W·m⁻² in the year 2100 without ever exceeding that value and thus assumes the imposition of emissions mitigation policies. The RCP 8.5 is based on the A2r scenario which combines assumptions about high population and relatively slow income growth with modest rates of technological change and energy intensity improvements, leading in the long-term to high energy demand and greenhouse gas emissions in the absence of climate change policies [22]. This set of data is the most up-to-date ensemble of high-resolution Regional Climate Model (RCM) projections. RCM historical data have been used for bias correction. The eight selected models have already been used over Burundi for climate projection [15] [23].

The data presented in **Table 1** are at 12 m height, the height at which IGEBU data were collected. The MERRA-2 and RCM wind speed data were however at 10 m height and the power law was used for data extrapolation at 12 m height [24].

2.3. Methods

The historical climate evaluation was conducted over the period 1980-2016 at hourly scale, seasonal scale and interannual scale. The diurnal analysis of the wind speed and wind direction was done from 7:00 AM to 6:00 PM while the nocturnal analysis was performed from 7:00 PM to 6:00 AM. Furthermore, the spatial distribution of WS and WPD has been adopted using Inverse Distance Weighted (IDW) interpolation method [25].

Table 2. Climatic models used.

Global Climate Model Name	Institute ID	Model Short Name	Model Origin Country
CanESM2	CCCma	CCCma	Canada
CNRM-CM5	CNRM-CERFACS	CNRM	France
EC-EARTH	ICHEC	ICHEC	Europe
IPSL-CM5A-MR	IPSL	IPSL	France
MIROC5	MIROC	MIROC	Japan
MPI-ESM-LR	MPI-M	MPI	Germany
NorESM1-M	NCC	NCC	Norway
GFDL-ESM2M	NOAA-GFDL	NOAA	USA

2.3.1. Bias Correction Method and Change Signal Detection

In this study, Empirical Statistical Downscaling (ESD) and Quantile Delta Mapping (QDM) methods were used for bias correction of RCM data. ESD took wind speed data from 0.44° resolution (~50 km × 50 km) to local climate, while QDM method served to bias adjust and detect changes. This computation was respectively implemented using ESD [26] and MBC package [27] in R program [28]. QDM preserves model-projected change signal in quantiles, while at the same time correcting systematic biases in quantiles of a modeled series with respect to observed values [29]. The conservation of changes follows directly from the quantile perturbation [30] and quantile delta change [31] methods, both of which apply simulated changes in quantiles overtop observed historical series [29].

2.3.2. Variability Analysis

The discrimination of years with high and low wind speeds was done using the standardized variable index (*I*) created by Mckee *et al.* [32] and which can be found in many articles [33] [34]:

$$I(i) = \frac{\phi_i - \overline{\phi}_m}{\sigma},\tag{1}$$

where ϕ_i , $\overline{\phi}_m$ and σ are respectively the value for the year *i*, the average and the standard deviation of the time series. Thus, in this paper, a year is considered normal if its index is included between -0.5 and +0.5. It is in high wind speeds if its index is greater than +0.5 and low wind speeds below -0.5. This interval remains criticized since it is relatively weak. But it clearly makes it possible to discriminate years with low wind speeds from years with high wind speeds. The standardized variable index also makes it possible to analyze the interannual variability of the key variable at considered spatial scales [15].

2.3.3. Wind Power Density

The wind speed data have been used to estimate the wind power, as follows [35]:

$$WP = \frac{1}{2}\rho A\phi^3, \tag{2}$$

Here WP is the power of the wind (in W), ρ is the air density (in kg·m⁻³), A is the swept area (in m²) and ϕ is the wind speed (in m·s⁻¹).

Therefore, the wind power density WPD (in W·m⁻²) is given by the following equation:

$$WPD = \frac{1}{2}\rho\phi^3,$$
(3)

2.3.4. Air Density

The air density ρ was computed as follows:

$$\rho = \frac{P}{RT},\tag{4}$$

where P is the atmospheric pressure (in mb), R is the universal gas constant

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 $(= 287 \text{ J} \cdot \text{kg}^{-1} \cdot \text{K}^{-1})$ and *T* is the air temperature (in K).

2.3.5. Turbulence Intensity

Wind turbulence is an important site characteristic, because high turbulence levels may decrease the power output and cause extreme loading on the wind turbine components. The turbulence intensity *TI* is defined as follows [36]:

$$TI = \frac{\sigma}{\phi_m},\tag{5}$$

where σ is the standard deviation and $\overline{\phi}_m$ is the average wind speed. In this paper, a *TI* value less than or equal to 0.10 is considered as indicating a low level of turbulence. *TI* value of 0.25 indicates moderate turbulence while a *TI* greater than 0.25 is considered as a high level of turbulence at the site.

2.3.6. Trend Analysis

For trend analysis, the Mann-Kendall's (MK) test (1945) [37] [38] was used at the threshold $\alpha = 0.05$. This trend test is worldwide used in meteorological variables trend detection [15] [39] [40] [41].

3. Results and Discussions

3.1. Wind Power Potential Analysis in the Historical Climate

3.1.1. MERRA-2 and RCM Performance

Figure 2 shows MERRA-2 performance to complement observed data and RCM performance to reproduce historical climate after bias correction. Indeed, the performance was measured with an objective method based on error function, which defines the goodness of fit, the Nash efficiency criterion, R^2 . Therefore,



Figure 2. MERRA-2 and bias corrected RCM performance.

 R^2 value of 0.95 was found between the ground data and MERRA-2 data used for each site. Furthermore, R^2 values of 0.95, 0.94, 0.92 and 0.95 were found at NHL, NIP, SHL and SIP, respectively between MERRA-2 and bias corrected RCM over the period 1980-2016.

3.1.2. Diurnal and Nocturnal Wind Speed and Wind Direction

Figure 3 presents average diurnal and nocturnal wind speed and wind direction at NHL over the period 1980-2016. The daywise average wind direction is East-South-East (ESE) with a frequency around 92%. The analysis revealed three parts of the wind rose where 25% of the winds blow with speeds below 4 $m \cdot s^{-1}$, 25% with speeds between 4 and 6 m·s⁻¹ while 42% have speeds between 6 and 7.7 $m \cdot s^{-1}$. The feature (NHL (day)) points out that during the day, the highest wind speed value is observed at 8 (or 2:00 PM) while the lowest wind speed value is observed at 1 (or 7:00 AM). Furthermore, the feature (NHL (day)) reveals the increase in the wind speeds from 1 to the 8th hour and the decline of the wind speeds from 8 to 12. On the other hand, nocturnal winds point three directions, South-East (SE), South-South-East (SSE) and South (S). From the S blow winds with speeds between 2 and 4 $m \cdot s^{-1}$ with a frequency of 32%. The wind blowing from the SSE has the speed between 2 and 4 $m \cdot s^{-1}$ with a frequency of 18%. Moreover, from the SE blows winds with a frequency of 50% where 80% and 20% of the winds respectively have speeds between 2 and 4 $m \cdot s^{-1}$, and speeds between 4 and 4.5 m·s⁻¹. The analysis of the feature (NHL(night)) reveals that the highest wind speed value at night is observed at 10 (or 4:00 AM) while from 3 to the 8th hour, we observe the lowest wind speeds. Comparing the day with the night, diurnal wind speeds are much higher than nightwise wind speeds. These findings show at which time of the day or night a given wind power turbine can expect receiving the maximum wind speeds.



Figure 3. Average wind speed and wind direction at NHL over the period 1980-2016.

Figure 4 presents the diurnal and nocturnal wind speed and wind direction at NIP. During the day, the wind rose points ESE and SE with the same frequency of 50%. The analysis shows that 50% of the ESE winds blow with speeds between 4 to 6 m·s⁻¹, while other 50% with speeds below 4 m·s⁻¹. On the other hand, 70% of SE winds have speeds between 6 and 7.6 m·s⁻¹ while 30% of SE winds blow with speeds below 6 m·s⁻¹. Furthermore, the feature (NIP (day)) shows the highest wind speed at 8 and the lowest value at 1. At night, the wind direction is SE where 90% of the winds blow with speeds below 4 m·s⁻¹ while 10% of the winds have speeds between 4 and 4.4 m·s⁻¹. The feature (NIP (night)) reveals the highest wind speed of the night at 10 (or 4:00 AM) with 4.4 m·s⁻¹.

Figure 5 presents the daywise and nightwise average wind speed and wind direction at SHL. The diurnal wind rose points ESE, SE and SSE with a frequency of 41% for both ESE and SE winds while SSE winds blow with a frequency of 18%. Furthermore, 62.5% of ESE winds blow with speeds below 4 m·s⁻¹ while 37.5% blow with speeds between 4 and 6 m·s⁻¹. Likewise, SE winds have all speeds between 6 and 7.7 m·s⁻¹. The SSE winds move with speeds between 4 and 6 m·s⁻¹ at SHL during the day. Moreover, the feature (SHL(day)) shows that the highest wind speed is observed at 8 (2:00 PM) while the lowest value is at 1 (7:00 AM). On the other hand, the nocturnal wind rose points SE, SSE and S with a frequency of 18%, 55% and 27%, respectively. Only 9% of the SSE winds have the speeds between 4 m·s⁻¹.

In the same way, **Figure 6** shows the diurnal and nocturnal average wind speed and wind direction at SIP over the period 1980-2016. The diurnal wind rose point SE, SSE and S as wind directions with a frequency of 41%, 33% and 26%, respectively. In fact, the analysis shows that 39% of SE winds blow with



Figure 4. Average wind speed and wind direction at NIP over the period 1980-2016.



Figure 5. Average wind speed and wind direction at SHL over the period 1980-2016.



Figure 6. Average wind speed and wind direction at SIP over the period 1980-2016.

speeds between 4 and 6 m·s⁻¹ while 61% blow with speeds below 4 m·s⁻¹. Furthermore, all SSE winds move with speeds between 6 and 7.9 m·s⁻¹ while 28% and 72% of S winds blow with speeds between 6 and 7.9 m·s⁻¹, and between 4 and 6 m·s⁻¹, respectively. Moreover, the feature (SIP(day)) shows that the highest wind speeds are observed at 7 (1:00 PM) and 8 (2:00 PM) while the lowest are at 1 (7:00 AM). On the other hand, the nocturnal wind rose provides SE, SSE and S as wind directions with a frequency of 19%, 25% and 56%, respectively. The SE and SSE winds blow with speeds below 4 m·s⁻¹ while S winds move with speeds between 4 and 5.3 m·s⁻¹. Indeed, the feature (SIP (night)) shows increase in wind speed from 6 to 10 where the highest wind speed is pointed at 10 (or 4:00 AM).

Overall, from the **Figures 3-6**, the findings revealed that at NHL, NIP, SHL and SIP, the average wind direction is between ESE and S. The highest wind speeds are observed in the afternoon before 4:00 PM, while at night the highest wind speeds are observed around 4:00 AM.

3.1.3. Seasonal Variation of Site Characteristic

Figure 7 presents the monthwise site characteristic including WS, SD, TI and AD at NHL, NIP, SHL and SIP over the historical period 1980-2016. SIP shows the highest seasonal WS. Furthermore, from May (5) to September (9), all the four sites present the highest WS of the year. This period coincides with the dry season in Burundi. During this season, hydropower plants experience the decrease of the water level in the reservoir which resulted in a difficult hydropower rationing. The figure also shows that the SD has high values during the rainy season between February and May, and in September and October while the lowest SD are observed in the dry seasons. The turbulence intensity (TI) is high (greater than 0.25) at NHL and NIP in March, April and October. All of the four sites reveal low TI values during the long dry season. Indeed, SIP does not reveal any month with high TI value, while only one month (October) of high TI is observed at SHL. In fact, low wind turbulence favors wind power generation while high wind turbulence may decrease it. Furthermore, the air density (AD) is higher at NIP and SIP than NHL and SHL. Indeed, this difference in AD is explained by the altitude between lowlands (NIP and SIP) and highlands (NHL



Figure 7. Monthly variation of average WS, SD, TI and AD at NHL, NIP, SHL and SIP.

and SHL). Air temperature and Pressure are higher across lowlands than highlands. In fact, low AD values decrease the wind power potential while higher values favor it. The feature also shows a slight increase in AD during the dry season all over the four sites.

3.1.4. Spatial Distribution of Monthly Wind Speed

Figure 8 presents monthwise wind speed from January to December at NHL, NIP, SHL and SIP over the period 1980-2016. In fact, from January to December, WS are higher in the south (SIP and SHL) than the north (NIP and NHL). The analysis shows that areas nearby the Lake Tanganyika (the south of NIP, SIP and the south-west of SHL) have higher WS than other parts. Furthermore, the year cycle can be divided into two main parts including October to April where WS are below $4.5 \text{ m}\cdot\text{s}^{-1}$ and from May to September where WS are higher than $4.5 \text{ m}\cdot\text{s}^{-1}$. Indeed, the findings revealed that the month of August has the highest WS of the year above $5.5 \text{ m}\cdot\text{s}^{-1}$.

3.1.5. Wind Speed Interannual Variability

Figure 9 presents the WS interannual variability at NHL, NIP, SHL and SIP over the period 1980-2016. The analysis of the standardized WS index reveals two main parts over the four sites including the period 1980-1996 in low wind speeds and the period 1997-2016 in high wind speeds. Indeed, Lawin *et al.* [15] attributed the cause of the difference between the two periods on environmental destruction. Furthermore, Havyarimana [42] showed negative impact of socio-political instability which started in Burundi in 1993 on the environment. His findings highlighted an increase in natural vegetation disturbance by anthropogenic activities over time.







Figure 9. Standardized WS index at NHL, NIP, SHL and SIP over the period 1980-2016.

Over the period 1980-1996, a further analysis of **Figure 9** revealed 9 years and 5 years of low WS at NHL and NIP, respectively, while 11 years of low WS were pointed out at SHL and SIP. The findings show that only NHL has 2 years of high WS over that period. Moreover, the period 1997-2016 presented 11 years, 8 years, 14 years and 9 years of high WS at NHL, NIP, SHL and SIP, respectively. Likewise, over the same period, NHL and NIP have 1 year of low WS, and SHL has 3 years, while SIP has no low WS over the period 1997-2016. Overall, the increase in years of high WS may have positive implication on wind power production, and hence reduce the electrical gap observed during the long dry season. Indeed, MK's test detected an upward trend in WS over the period 1980-2016 at NHL, NIP, SHL and SIP. In fact, at significance level equals to 0.05, the Kendall's tau of 0.31, 0.38, 0.36 and 0.49 with their respective p-values equal to 0.0098, 0.0016, 0.0033 and < 0.0001 were revealed at NHL, NIP, SHL and SIP, respectively. This increasing trend in wind speed is consistent with the findings in many studies from the East Africa [15] [43].

3.1.6. Wind Power Density over the Baseline Period

Figure 10 presents the monthly distribution of wind power density (WPD) over the reference period 1981-2010. The figure highlights areas in low and high WPD for the twelve months of the year. Therefore, from June to September, we have months in high WPD with values greater than 60 W·m⁻² while from November to March, we have months in low WPD with values less than 25 W·m⁻². The findings show that the months of April, May and October have WPD between 25 and 60 W·m⁻². Furthermore, within the four sites, SIP is revealed to have higher WPD than other sites over the entire baseline period with values greater than 120 W·m⁻² during the summer time. Indeed, from the results, the four sites are only promising for installing wind turbine for small-scale power



Figure 10. Spatial distribution of average WPD at NHL, NIP, SHL and SIP over the period 1981-2010.

generation. Otherwise, large-scale wind power generation require that mean wind speeds exceed 7 $\text{m}\cdot\text{s}^{-1}$ to generate at least 200 W·m⁻² [24].

According to the ISTEEBU report [44], Burundi hydroelectricity production is characterized by a lot of fluctuation due to its dependence on rainfall; which impels the importation of hydroelectricity and the usage of fossil fuel for electricity generation especially during the dry season. For example, in 2007, the report shows that 39% of electricity was imported from RUZIZI I and RUZIZI II of the Democratic Republic of the Congo. In 2013, Burundi started its three electricity generators using fossil fuel to generate 14% (23 MW) of national electricity production. The usage of the power generators reduced importation to 37%. However, the cost of 1 kWh for consumers was increased up to 47% compared to the cost of 2007. Therefore, wind power can contribute to reduce electricity importation and the usage of fossil fuel for electricity generation. Moreover, the demand of water for irrigation is seasonal and it is found that the highest wind power density coincides with the dry season which may favor wind powered water pumping applications.

3.2. Future Wind Power Potential

We recall that the future climate has been spatially and temporally analyzed using a multi-model ensemble of eight Regional Climate Models (CCCma, CNRM, ICHEC, IPSL, MIROC, MPI, NCC and NOAA) over the periods 2019-2040 and 2071-2100 according to the reference period 1981-2010. Specifically, a multi-model mean has been computed and adopted for each emission scenario (RCP 4.5 and RCP 8.5).

3.2.1. Projected Wind Speed over the Period 2019-2040

Figure 11 and Figure 12 present the spatial projected monthly WS following RCP 4.5 and RCP 8.5, respectively over the period 2019-2040 according to the baseline period 1981-2010. The analysis of Figure 11 shows that from October to March, projected WS are between 3 and 5 $m \cdot s^{-1}$ all over the study area. Indeed, the months April-September are expected to experience WS above 5 m·s⁻¹. The months June-August are forecasted to have the highest WS above 6.5 $m \cdot s^{-1}$. Among the four sites, SIP is expected to experience the highest WS. Furthermore, the south of NIP will experience higher WS than other parts of NIP. Moreover, the southern parts of SHL are projected to be in higher WS than other parts of SHL. Likewise, the analysis of Figure 12 reveals that from October to April, the WS are projected to be between 3.5 and 5.2 $\text{m}\cdot\text{s}^{-1}$ where the month of April and March (the south) will experience WS between 4.5 and 5.2 m·s⁻¹. Furthermore, the months May-September are forecasted to experience WS between 6 and 8 $m \cdot s^{-1}$. Higher WS are projected at SIP than other parts over the period 2019-2040 following RCP 8.5. Comparing findings from RCP 4.5 with RCP 8.5, WS values are higher in RCP 8.5 with a slight difference except at SIP in September.

3.2.2. Projected Wind Speed over the Period 2071-2100

Figure 13 and **Figure 14** present the spatial forecasted monthly WS following respectively RCP 4.5 and RCP 8.5 over the period 2071-2100 according to the reference period 1981-2010. The analysis of **Figure 13** shows that the highest WS of the year cycle are expected in May-September with values between 6.5 and 8.5 m·s⁻¹ while the lowest WS are projected in October-April with values between 4.5 and 6.5 m·s⁻¹. Furthermore, SIP is forecasted to have higher WS



Figure 11. Monthly projected average WS at NHL, NIP, SHL and SIP following RCP 4.5.



Figure 12. Monthly projected average WS at NHL, NIP, SHL and SIP following RCP 8.5.



Figure 13. Monthly projected average WS at NHL, NIP, SHL and SIP following RCP 4.5.

than other parts from January to December. The south of NIP and the South-West of SHL will experience higher WS than other parts of NIP and SHL, respectively.

The analysis of **Figure 14** reveals that from October to April, the WS are projected to vary between 4.8 and 6.5 $\text{m}\cdot\text{s}^{-1}$ while the months May-September are projected to experience WS between 6.5 and 10 $\text{m}\cdot\text{s}^{-1}$. Higher WS are projected at SIP than other parts of the study area. Comparing the findings from RCP 4.5 with RCP 8.5 over the period 2071-2100, WS values are higher in RCP



Figure 14. Monthly projected average WS at NHL, NIP, SHL and SIP following 8.5.

8.5 with a slight difference except at SIP in September.

Overall, comparing the projection over the period 2019-2040 with the projections over the period 2071-2100 according to the baseline period 1981-2010, the findings show that projected WS are higher over the far future than near future. Indeed, higher projected wind speeds will obviously have positive impact on wind power production.

3.2.3. Interannual Projected Wind Speed

Figure 15 presents interannual projected WS over the period 2019-2040 following RCP 4.5 (RCP 4.5 (a) and RCP 8.5 (a)), and over the period 2071-2100 following RCP 4.5 (RCP 4.5 (b) and RCP 8.5 (b)) with their reference period 1981-2010 (obs). The figure shows that future interannual WS will be characterized by a lot of fluctuations and a general increase in WS. Furthermore, the features show that WS values projected under RCP 4.5 and RCP 8.5 are very close all over the four sites. The analysis revealed that over the period 2019-2040, both scenarios projected an average increase in interannual WS of 0.9 and 0.8 $m \cdot s^{-1}$ in the highlands and lowlands, respectively.

Likewise, over the period 2071-2100, an average increase of 2.1 and 2 m·s⁻¹ is respectively expected across the highlands and lowlands in RCP 4.5 while in RCP 8.5, an average increase of 2.4 and 2.3 m·s⁻¹ are projected in the latter respective sites. However, in both climate scenarios, all interannual average wind speeds will hover between 4 and 8 m·s⁻¹. Indeed, over the period 2071-2100, projected WS are expected to be slightly higher in RCP 8.5 than RCP 4.5. Overall, the findings obviously shows that projected wind power are expected to increase. These findings are consistent to other studies all over the world [2] [4] [43] [45].



Figure 15. Observed and interannual projected average WS at NHL, NIP, SHL and SIP.

3.2.4. Wind Power Density Projected Changes

Figure 16 and **Figure 17** show monthwise and annual projected changes in WPD over the periods 2019-2040 and 2071-2100 at NHL, NIP, SHL and SIP according to the baseline period 1981-2010. The figures used violins to illustrate WPD projected changes. The red dot shows the mean of projected WPD changes. The thicker part indicates that the monthly WPD change value in that section of the violin has higher frequency, and the thinner part implies lower frequency. The bottom of the violin shows the minimum WPD change value while the top of the violin means the maximum value. For more clarity, only months corresponding to the maximum and minimum changes have been plotted with the violin. Furthermore, the first crossbar of the violin marks the 1st quartile, the second crossbar is the 2nd quartile (median) and the last is the 3rd quartile.

Therefore, over the period 2019-2040 following RCP 4.5, an average WPD change of 28.04 W·m⁻², 30.27 W·m⁻², 28.15 W·m⁻² and 31.1 W·m⁻² is projected at NHL, NIP, SHL and SIP, respectively. Although these changes are close, the analysis shows that WPD changes in quartiles for lowlands (NIP and SIP) are higher than highlands (NHL and SHL). All over the four sites, the month of August is projected to experience the highest WPD changes while January will have the lowest. In August the highest WPD change value is expected at SIP with 50.6 W·m⁻², while in January the lowest change will be at NHL with 12.5 W·m⁻². Likewise, following RCP 8.5, the month of September is expected to have the highest WPD changes all over the study area. Indeed, in the lowlands, September shows the higher changes than highlands where SIP is expected to experience the highest WPD change value equal to 99.6 W·m⁻². Overall, in RCP 8.5, NHL, NIP, SHL and SIP are projected to experience WPD changes equal to 28.3 W·m⁻²,



Figure 16. Projected monthly and annual average WPD changes over the period 2019-2040.



Figure 17. Projected monthly and annual average WPD changes over the period 2071-2100.

31.24 $W \cdot m^{-2}$, 28.42 $W \cdot m^{-2}$ and 47.35 $W \cdot m^{-2}$, respectively. Both the scenarios agree on January as a projected month to experience the lowest WPD changes across the lowlands and highlands.

On the other hand, over the period 2071-2100 (**Figure 17**) following RCP 4.5, averages WDP change equal to 83.96 W·m⁻², 94.17 W·m⁻², 85 W·m⁻² and 116.16 W·m⁻² are projected at NHL, NIP, SHL and SIP, respectively. The analysis shows that WPD changes in lowlands are higher than highlands. In addition, within the

two regions (lowlands and highlands), changes are higher in the south than north. Indeed, all over the four sites, September is projected to experience the highest WPD changes while January will have the lowest. In January the lowest WPD change value is forecasted at NHL with 37.8 W·m⁻², while in September the highest WPD change is expected at SIP with 198.8 W·m⁻². In the same way, following RCP 8.5 over the same period, the month of September is expected to have the highest WPD changes for all the four sites. Indeed, September shows the highest WPD change equal to 259.8 W·m⁻² at SIP while January projects the lowest WPD change equal to 48.7 W·m⁻² at NHL.

Overall, NHL, NIP, SHL and SIP are expected to experience WPD changes equal to 103.82 $W \cdot m^{-2}$, 116.9 $W \cdot m^{-2}$, 104.51 $W \cdot m^{-2}$ and 152.39 $W \cdot m^{-2}$, respectively. Comparing the findings from the two scenarios over the period 2071-2100, both agree on January and September as months projected to experience the lowest and the highest WPD changes, respectively throughout the lowlands and highlands.

Furthermore, comparing the period 2019-2040 with 2071-2100, both of them agree on the increase in the projected wind power density according to the reference period 1981-2010. In fact, no negative WPD change was found.

3.2.5. Spatial Distribution of Wind Power Density Changes

Figure 18 presents the spatial distribution of annual WPD changes over the period 2019-2040 (19/40) and over the period 2071-2100 (71/100) following RCP 4.5 and RCP 8.5 according to the reference period 1981-2010. The figure shows that changes are expected to be higher in RCP 8.5 than RCP 4.5 over both periods. In fact, changes in RCP 4.5 will be less than 50 W·m⁻² while changes in RCP 8.5 will be higher than 75 W·m⁻².

Moreover, the figure highlights that changes are projected to increase over the two time periods within each climate scenario. Furthermore, western areas nearby the Lake Tanganyika are expected to experience higher increase in WPD



Figure 18. Annual spatial distribution of projected changes in wind power density.

than other areas. Indeed, the entire SIP, the south of NIP, the south and the south-west of SHL are projected to experience higher changes in wind power density than other parts of the study area.

3.2.6. Future Wind Speed Potential on Different Heights in the Surface Layer

The axes of modern wind turbines are located at different heights in layer up to 165 m thick [35]. For most of them, the cut-in speed is $2.5 - 5 \text{ m} \cdot \text{s}^{-1}$, the rated speed is $10 - 15 \text{ m} \cdot \text{s}^{-1}$, the cut-out speed is within $24 - 25 \text{ m} \cdot \text{s}^{-1}$ [36]. Indeed, all these values refer to wind speed at the height of axis of the wind turbine. Recall that in surface layer, the wind speed increases with increasing height above the ground [24]. Therefore, it is necessary to take into consideration that the wind speed at hub height has higher values. Figures 19-21 show projected wind speed at 50 m, 100 m and 150 m heights. The analysis revealed that from 12 m to 50 m height, the lowest expected increase in wind speed over the four sites is $1.1 \text{ m} \cdot \text{s}^{-1}$, while the highest increase of $2.2 \text{ m} \cdot \text{s}^{-1}$ and $2.8 \text{ m} \cdot \text{s}^{-1}$ are projected at 150 m height for the four sites over the period 2019-2040 and 2071-2100, respectively under both RCP 4.5 and RCP 8.5. Overall, 150 m may be considered as the best height for wind power generation. This can help to reduce electricity importation and the fossil fuel dependence to generate electricity and the cost of 1 kWh especially during the dry season.

The findings from this study projecting increase in WS which implies increase in WPD are consistent with many studies from different regions of the globe as well as studies from East Africa [2] [45] [46] [47] [48]. Indeed, Lawin *et al.* [15] reported increasing trend in wind speed at Northern Imbo plain while Collins *et al.* [14], predicted that Burundi would experience more extreme weather conditions in the future characterized by high winds; which is consistent with our findings.



Figure 19. Spatial distribution of projected annual wind potential at 50 m.



Figure 20. Spatial distribution of projected annual wind potential at 100 m.



Figure 21. Spatial distribution of projected annual wind potential at 150 m.

4. Conclusions

This study analyzed the impact of climate change on wind power potential over the periods 2019-2040 and 2071-2100 using two emission scenarios RCP 4.5 and RCP 8.5 according to the reference period 1981-2010.

The wind over the historical period 1980-2016 has been analyzed at hourly scales, seasonal scales and interannual scales. At hourly scales, diurnal wind direction was hovering between South-East and South, and the wind speeds were revealed higher in the afternoon. At monthly-scales, the month of August was revealed to have the highest WS of the year with values above $5.5 \text{ m}\cdot\text{s}^{-1}$. Furthermore, the analysis of the standardized WS index revealed two main parts over the four sites including the period 1980-1996 in low wind speeds and the period 1997-2016 in high wind speeds. Indeed, MK's test detected an upward trend in WS over the period 1980-2016. Moreover, the findings showed that the months June-September are in high WPD with values greater than 60 W·m⁻² while the months November-March are in low WPD with values less than 25

 $W \cdot m^{-2}$. Indeed, within the four sites, SIP was revealed to have higher WPD than other sites over the entire baseline period with values greater than 120 $W \cdot m^{-2}$ during the summer time.

In the future climate, monthly WS are expected to increase all over the studied sites in both climate scenarios over the periods 2019-2040 and 2071-2100 which will obviously increase WPD. Therefore, following RCP 4.5 over the period 2019-2040, an average WPD change of 28.04 W·m⁻², 30.27 W·m⁻², 28.15 W·m⁻² and 31.1 W·m⁻² is projected at NHL, NIP, SHL and SIP, respectively. All over the four sites, the month of August is projected to experience the highest WPD change. On the other hand, following RCP 8.5, the month of September is expected to have the highest WPD change all over the study area. The sites NHL, NIP, SHL and SIP are projected to experience WPD changes equal to 28.3 W·m⁻², 31.24 W·m⁻², 28.42 W·m⁻² and 47.35 W·m⁻², respectively in RCP 8.5. Likewise, over the period 2071-2100 following RCP 4.5, an average WDP change of 83.96 W·m⁻², 94.17 W·m⁻², 85 W·m⁻² and 116.16 W·m⁻² is respectively projected at NHL, NIP, SHL and SIP. On the other hand, following RCP 8.5, the sites NHL, NIP, SHL and SIP are expected to experience WPD change of 103.82 W·m⁻², 116.9 W·m⁻², 104.51 W·m⁻² and 152.39 W·m⁻², respectively. Both the scenarios agree on January and September as months expected to experience the lowest and the highest WPD change, respectively over the studied area. They also agreed on SIP as the site expected to experience the highest WPD. The findings highlighted that areas nearby the Lake Tanganyika are expected to experience higher increase in WPD than other areas.

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Conflicts of Interest

The authors declare no conflicts of interest regarding the publication of this paper.

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